# GEOL 414/514

# **ACIDS AND BASES**

# **Chapter 5**

# LANGMUIR

# SIGNIFICANCE & MEASUREMENT OF pH

 $H_2O \leftrightarrow H^+ + OH^-$ 

$$K_w = \frac{[H^+][OH^-]}{[H_2O]} = 10^{-14.00}$$

pH = -log [H+]

Most reactions involve or are controlled by pH:

- 1. Aqueous acid-base equilibria (hydrolysis, polymeriz'n
- 2. Adsorption, H<sup>+</sup> & OH<sup>-</sup> compete with other ions
- 3. Formation of metal-ligand complexes (H\*&OH<sup>-</sup> compet'n)
- 4. Oxidation-reduction rea'ns (H<sup>+</sup> again)
- 5. Solubility & rate of mineral dissolution

# SIGNIFICANCE & MEASUREMENT OF pH

• Ultimate reference for pH & Eh measurements is H<sub>2</sub> electrode - bubble H<sub>2</sub> over Pt electrode

$$\frac{1}{2}H_2(g) = H^+ + e^-$$

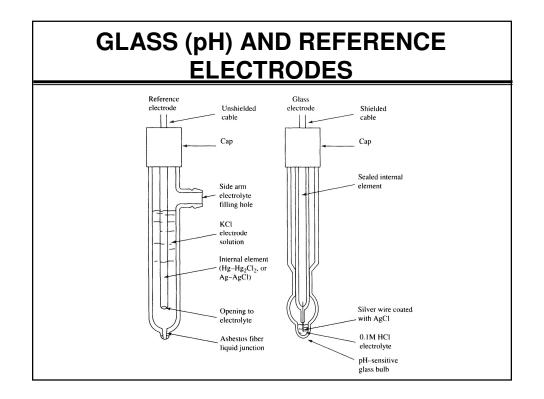
 Cannot conveniently use the H<sub>2</sub> for routine measurements; use a glass electrode (pH) with a reference electrode - usually a Ag-AgCl or Hg-Hg<sub>2</sub>Cl<sub>2</sub> cell. Using a buffer solution, pH determinations are made

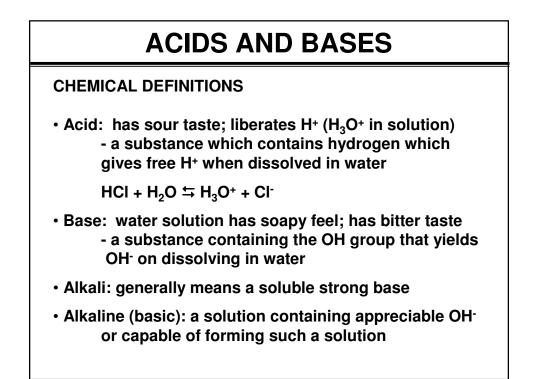
# SIGNIFICANCE & MEASUREMENT OF pH

The pH of the unknown is obtained from:

$$pH_x = pH_b + \frac{(E_x - E_b)}{1.984 \times 10^{-4} \text{ T}}$$

- The electrodes measure voltages which the pH meter converts and displays as pH units
- The most important function of the reference electrode is to provide known and stable reference voltage
- There is a liquid junction potential due to the nature of the reference electrode - may be important in solutions of very low *I*





# **ACIDS AND BASES - 2**

### **GEOLOGIC USAGE**

- The terms acid, base & alkali have wide and varied meanings in geology
- Many oxides of nonmetals dissolve in H<sub>2</sub>O to form acids
- Term "acid oxide" used for nonmetallic oxide regardless of solubility (SiO<sub>2</sub> & rocks with high % SiO<sub>2</sub>)
- Term "basic oxide" used for any metal oxide
- Text (& this course) uses strict chemical definitions
- Quartz-rich rock is felsic or silic
- So-called basic rocks are mafic
- Rocks rich in Na or K are called alkalic or alkaline

# ACIDS AND BASES - 3

THE pH

Neutralization reactions

 $HCI + NaOH \rightarrow HOH + NaCI$ 

More strictly:

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H^+ + CI^- + Na^+ + OH^- \rightarrow H_2O + Na^+ + CI^-
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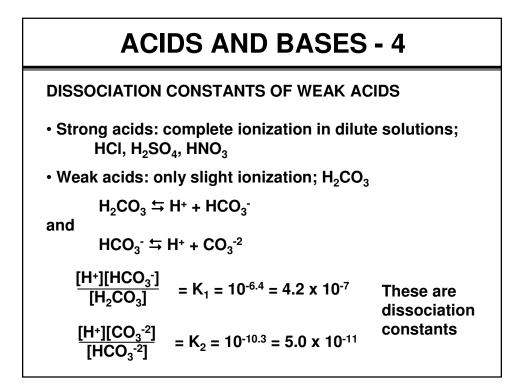
• More simply, since Na<sup>+</sup> & Cl<sup>-</sup> are unaffected:

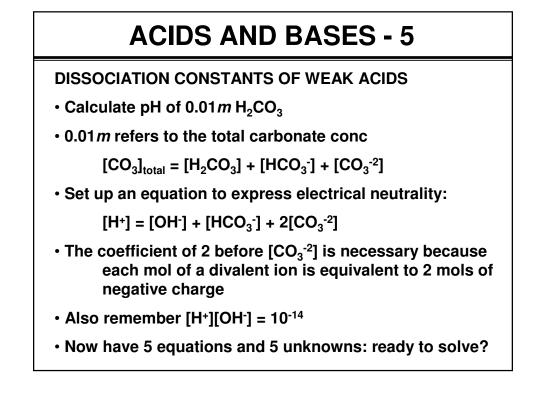
 $OH^{-} + H^{+} \rightarrow H_{2}O$ 

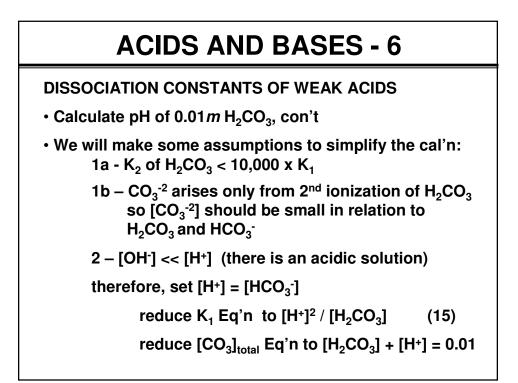
• Dissociation of  $H_2O$ :  $H_2O \leftrightarrows H^+ + OH^-$ 

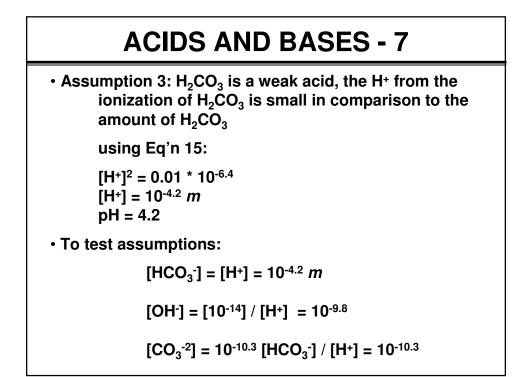
K<sub>H2O</sub> = [H<sup>+</sup>][OH<sup>-</sup>] = 10<sup>-14</sup>

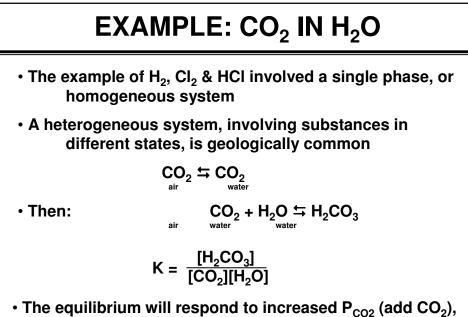
• The product, [H+][OH<sup>-</sup>], is constant for ALL H<sub>2</sub>O solutions











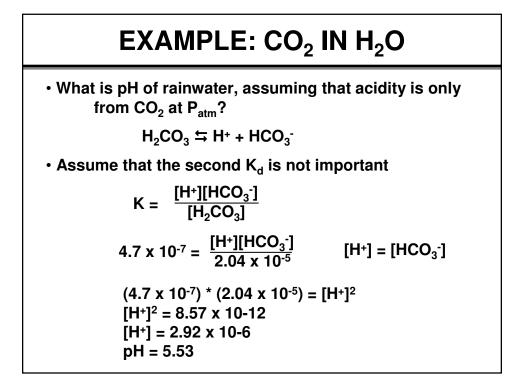
decreased P<sub>CO2</sub> (heat, create vacuum)

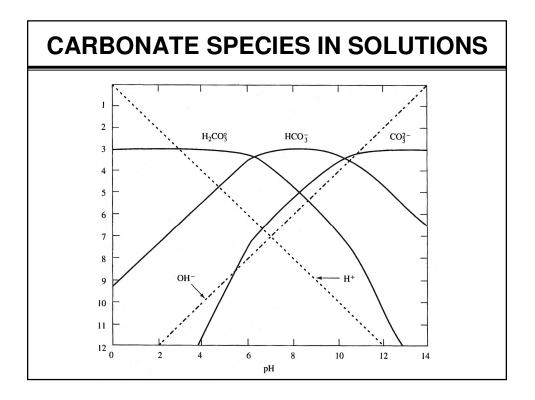
# • Solubility of CO<sub>2</sub> = 0.76 L/L H<sub>2</sub>0 at t = 25 °C, P<sub>CO2</sub> = 1 bar • 0.75 L CO<sub>2</sub>; 0.76L/24.5 L/mole = 0.31 moles CO<sub>2</sub>

- [Notice that H<sub>2</sub>O = 1000/18.016 = 55.5 *m*]
- For CO<sub>2</sub> above: K =  $[H_2CO_3]/[CO_2] = 0.031/1 = .031 \approx 10^{-1.5}$
- How much H<sub>2</sub>CO<sub>3</sub> in water exposed to ordinary air?
- Air contains 0.03% CO2 by volume; volume fraction = 0.0003 & partial pressure = 0.00003 bar

 $K = 0.031 = [H_2CO_3] / 0.0003$ 

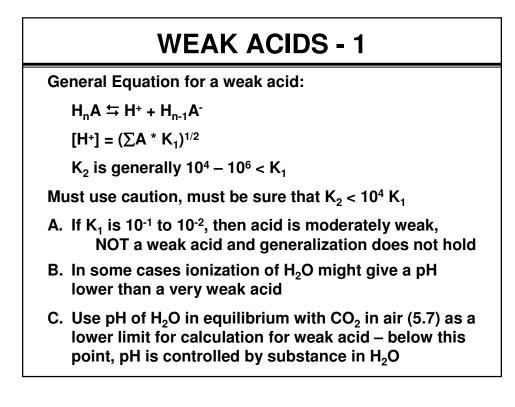
 $[H_2CO_3] = 0.031 \times 0.0003 = 10^{-1.5} \times 10^{-3.5} = 10^{-5} m$ 

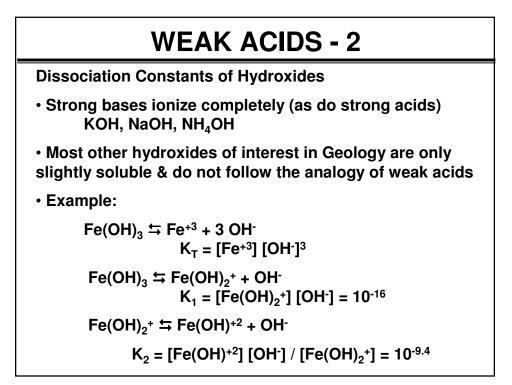




# **ESTIMATING IONIC CONCENTRATIONS**

- It is very often desirable to determine the ionic conc and ion distribution in natural systems
- We can use dissociation and hydrolyses equations, and associated K<sub>eq</sub>, to determine the major, minor and dominant species
- See the example of H<sub>2</sub>CO<sub>3</sub> in the text
- When pH < 6.4 then H<sub>2</sub>CO<sub>3</sub> dominates
- When pH = 6.4 10.3 then  $HCO_3^{-1}$  dominates
- When pH> 10.3 then CO<sub>3</sub><sup>-2</sup> dominates
- The above is true for these pH ranges regardless of concentrations in solution & whether or not any other solutes are present

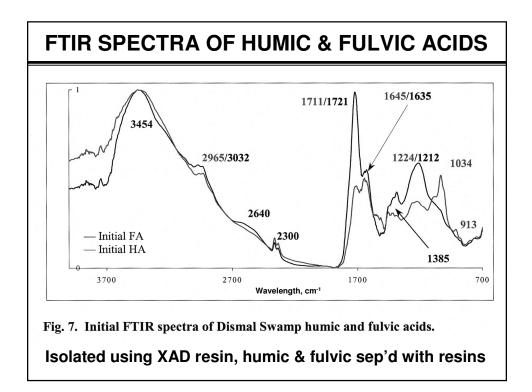


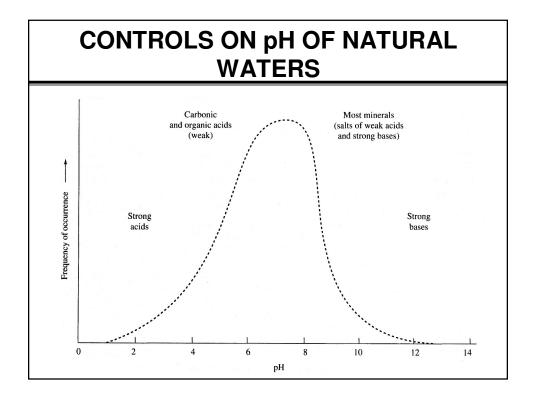


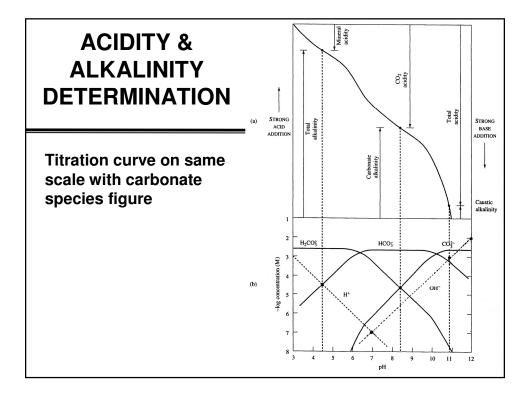
# WEAK ACIDS - 3FeOH+2 \$\IPS\$ Fe+3 + OH: $K_3 = [Fe^{+3}] [OH^{-1}] = 10^{-11.7}$ $K_3 = [Fe^{+3}] [OH^{-1}] = 10^{-11.7}$ $Fe(OH)_3 + OH^{-1} \IPS Fe(OH)_4^{-1}] [OH^{-1}]$ $K_A = [Fe(OH)_4^{-1}] [OH^{-1}]$ $Fe(OH)_3 \IPS Fe(OH)_3^{\circ}$ $K_{aq} = [Fe(OH)_3^{\circ}] = 10^{-7.0}$ In a natural system, there may also be $Fe(OH)_2$ present; ionization and speciation of this form would also have to be considered

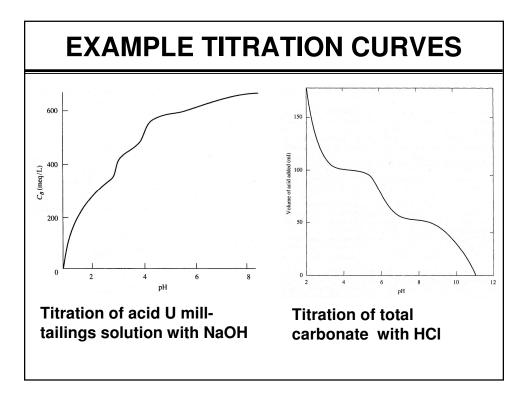
# HUMIC AND FULVIC ACIDS

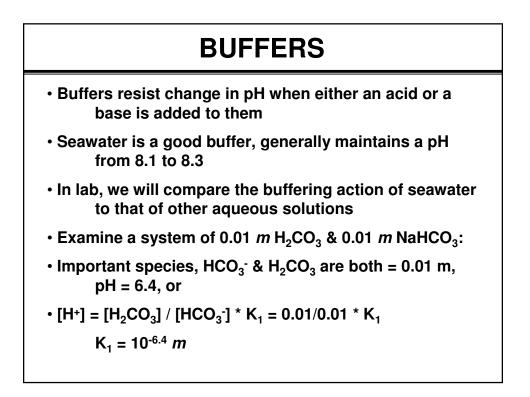
- Soil humic substances (60-80% SOM): fulvic acid - low M.W., sol in acid or base humic acid - med M.W., sol in base, insol acid humin - highest M.W., insol in acid or base
- Soil nonhumic substances: carbohydrates, proteins, peptides, amino acids, fats, waxes
- Humic substances are acidic, dark-colored, partially aromatic, M.W.'s range from 100's -1000's g FW<sup>-1</sup>
- Aquatic humic substances:
  - polyelectrolytic acids that can be isolated from water by sorption onto XAD or weak-base exch resin;
    - M.W., 500-5000 g/mol
    - humic & fulvic acids









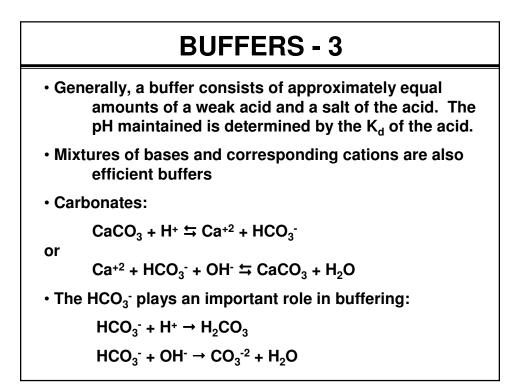


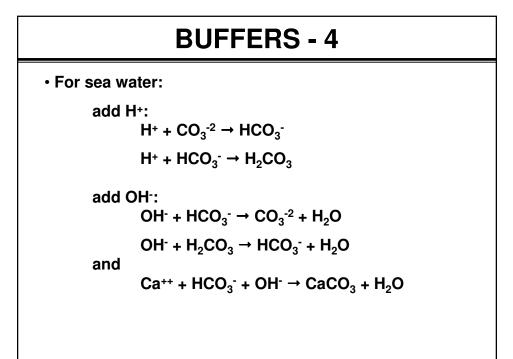
# **BUFFERS - 1**

H<sub>2</sub>CO<sub>3</sub> & NaHCO<sub>3</sub> example continued:

1. Add acid:  $H^+ + HCO_3^- \rightarrow H_2CO_3$ 

- the ratio of [H<sub>2</sub>CO<sub>3</sub>] : [HCO<sub>3</sub>-] changes which does change the [H<sup>+</sup>] but not very much unless the ratio is greatly changed;
- change ratio from 1 to 2, pH changes from 6.4 to 6.1
- Add base:  $OH^- + H_2CO_3 \rightarrow HCO_3^- + H_2O$
- change ratio in opposite direction, pH changes from 6.4 to 6.7





# ELECTROLYTE SOLUTIONS & ION-ION ASSOCIATIONS

- In electrolytic solutions, some ions & molecules form combinations, called complexes. Many of these are commonly referred to as ion pairs.
- They behave as a unit and are treated as a thermodynamic entity.

**Chloride complexes** 

 Formation is common with many types of cations, especially transition metals, over a wide range of geochemical conditions

CdCl<sup>+</sup>, PbCl<sup>+</sup>, ZnCl<sup>+</sup>, ZnCl<sub>3</sub><sup>-</sup>

Ion sorption, mobility, volatility affected